

# Natural iron fertilization controls deep-water carbon and opal flux in the Southern Ocean

**Mike Lucas (UCT), Raymond Pollard (NOC)  
& the Crozet team**

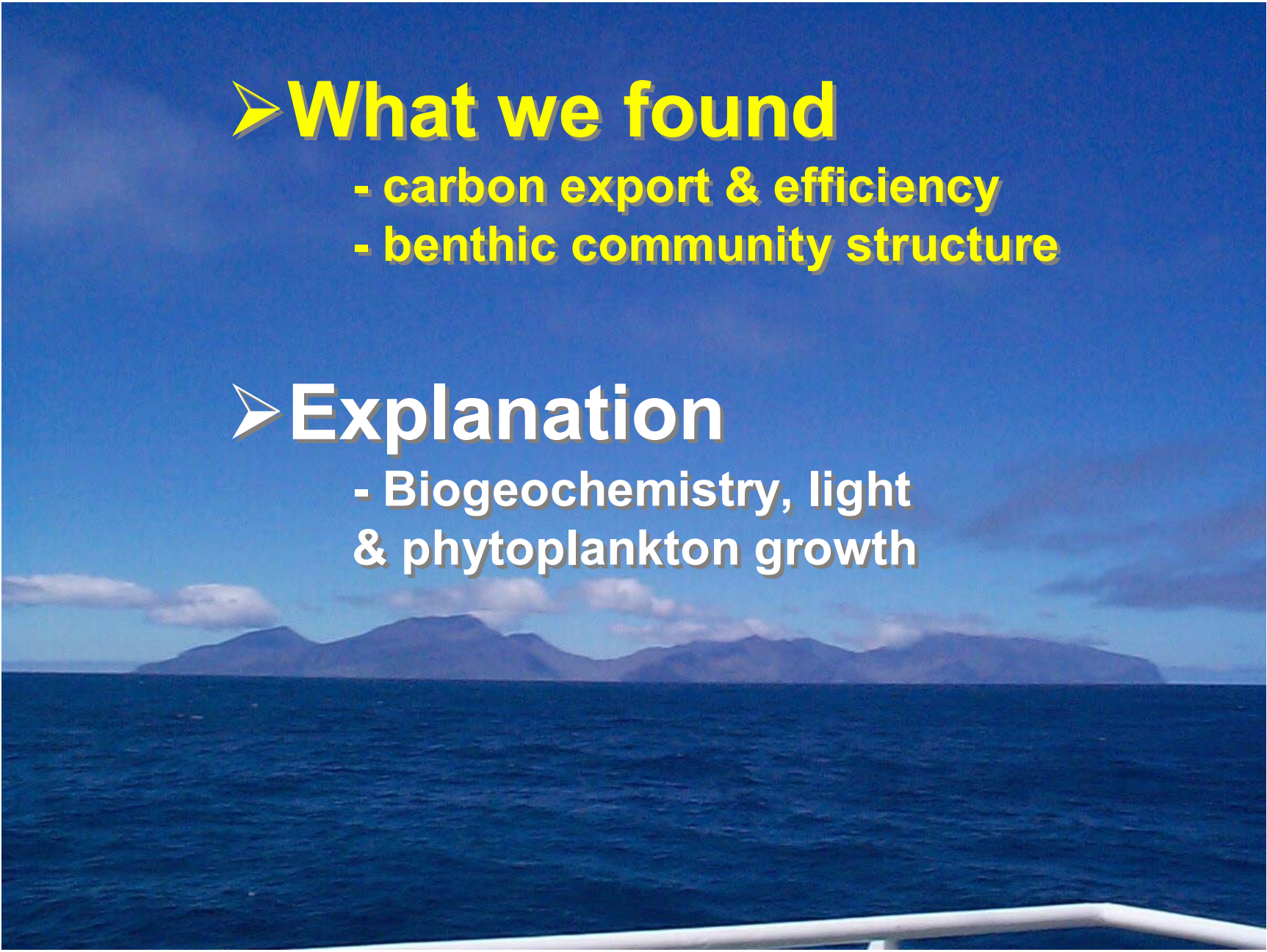
DSR II, 54: 2007; Nature, in revision

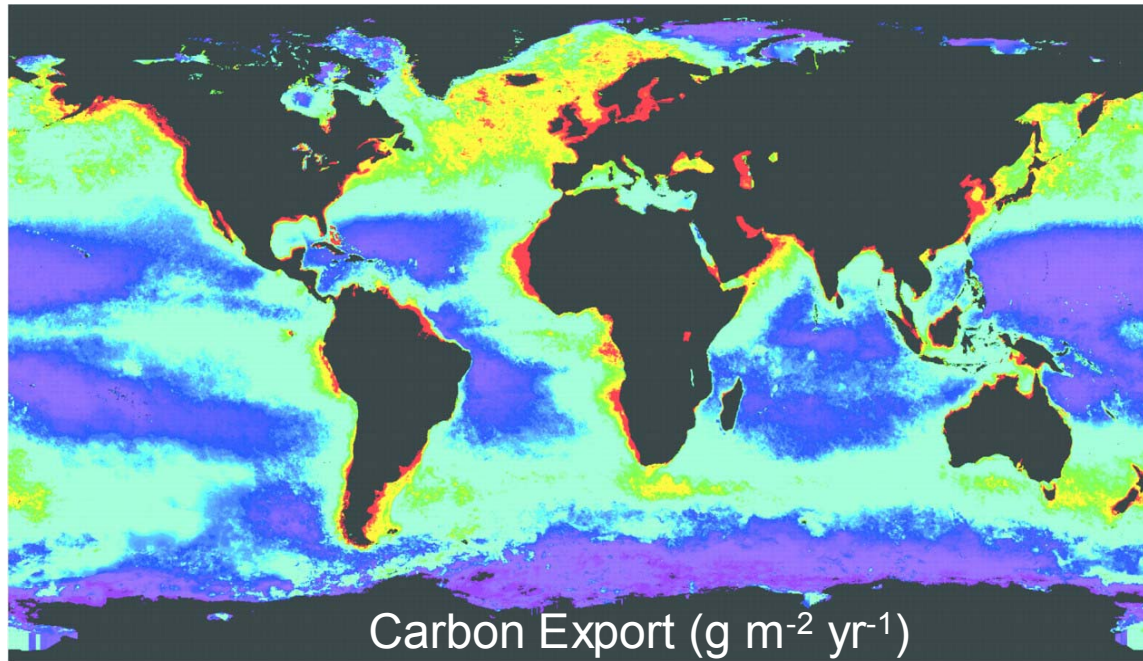
## ➤ **What we found**

- carbon export & efficiency
- benthic community structure

## ➤ **Explanation**

- Biogeochemistry, light & phytoplankton growth



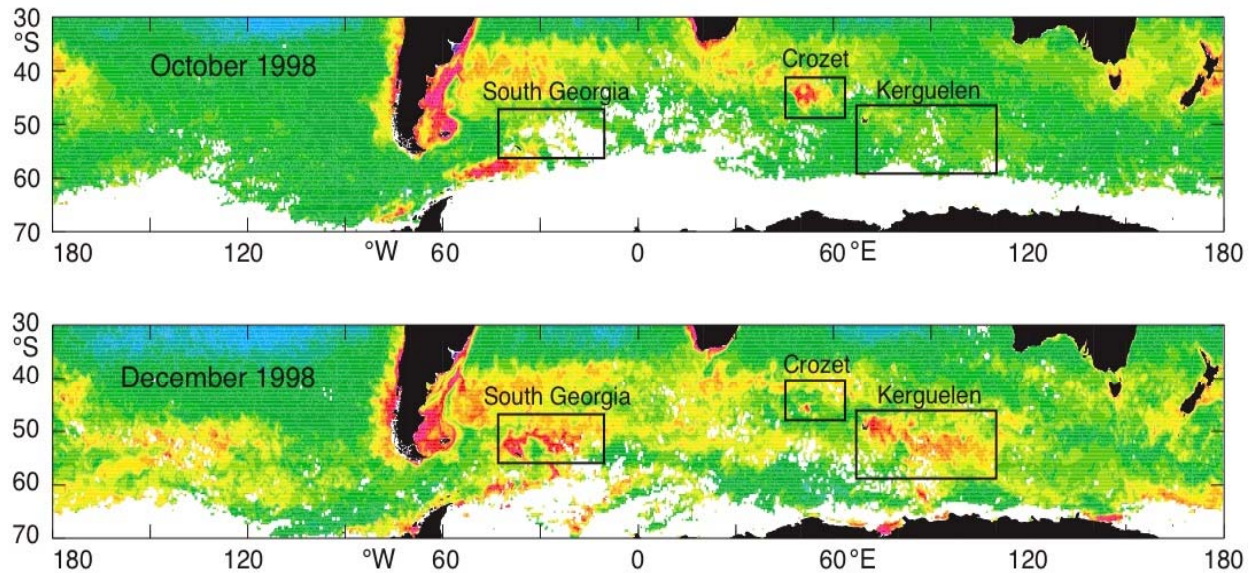


**Low C. export in high nutrient Southern Ocean.**

**Due primarily to Fe limitation.**

**Would alleviation of Fe limitation enhance C. export?**

# Test this hypothesis from observations of blooms around sub-Antarctic Islands

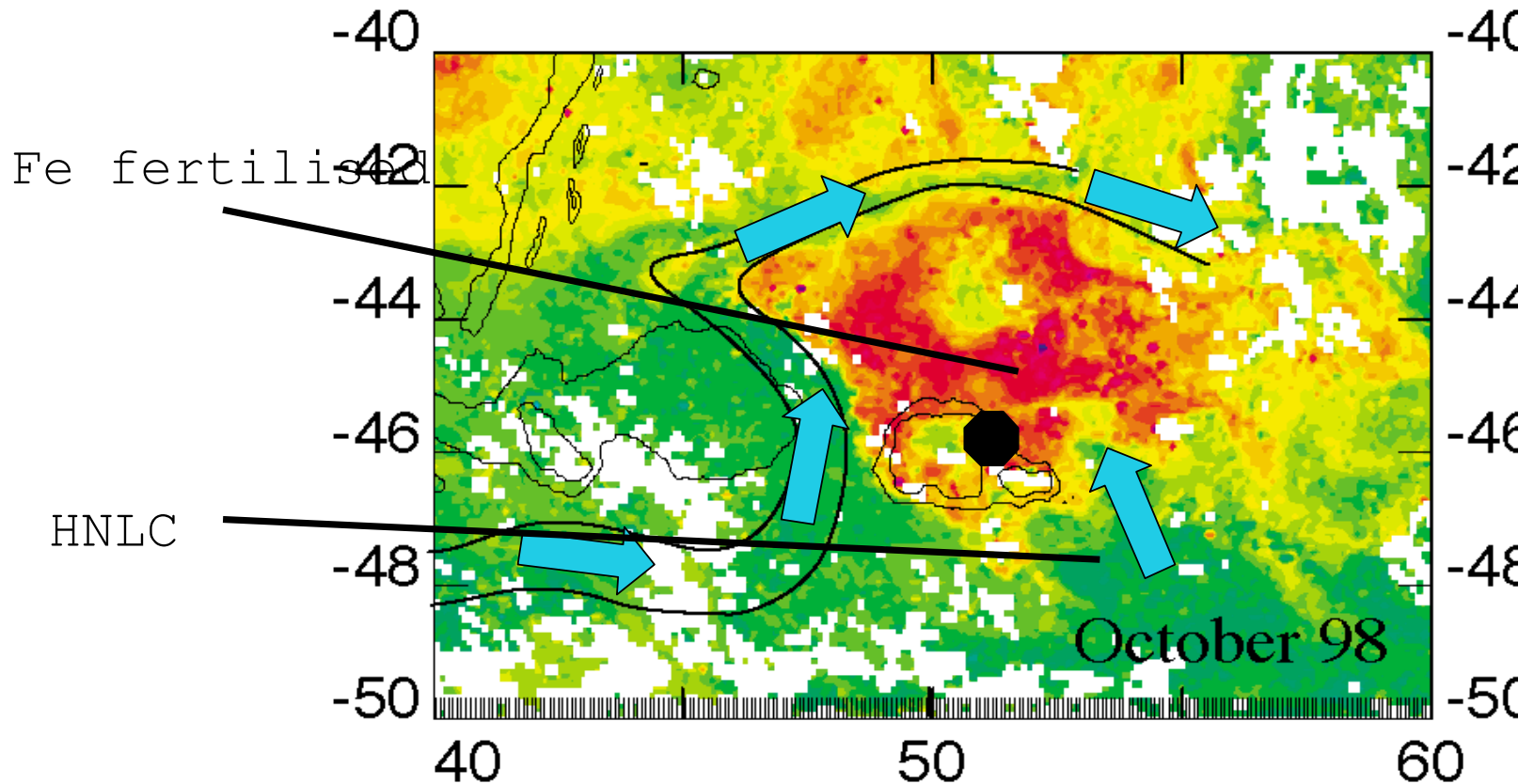


Pollard et al., DSR II, 54: 2007

Fig. 1

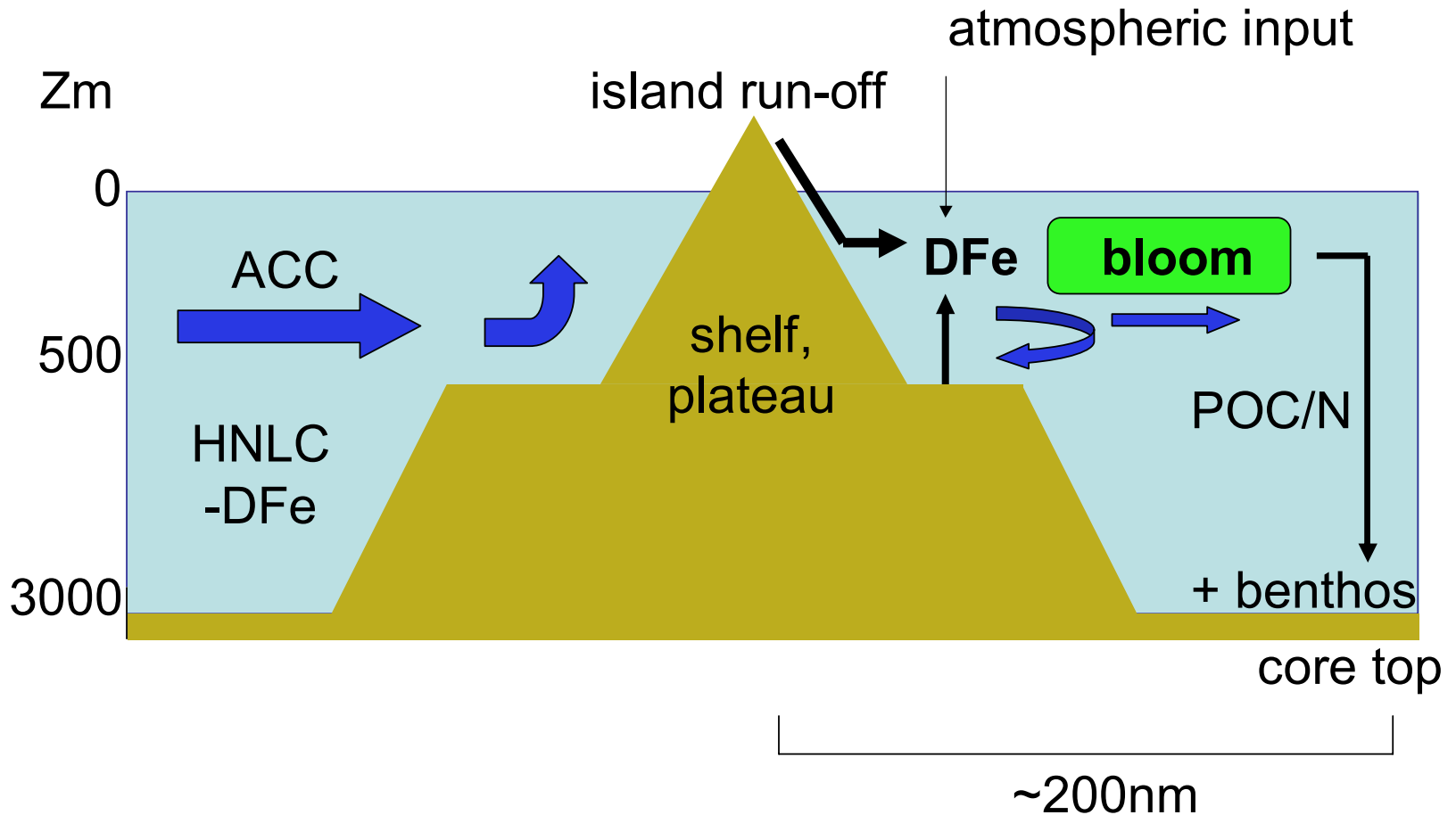
# Test carbon export wrt Fe Fertilisation

(Pollard, Sanders, Lucas, Statham et al. 2007)



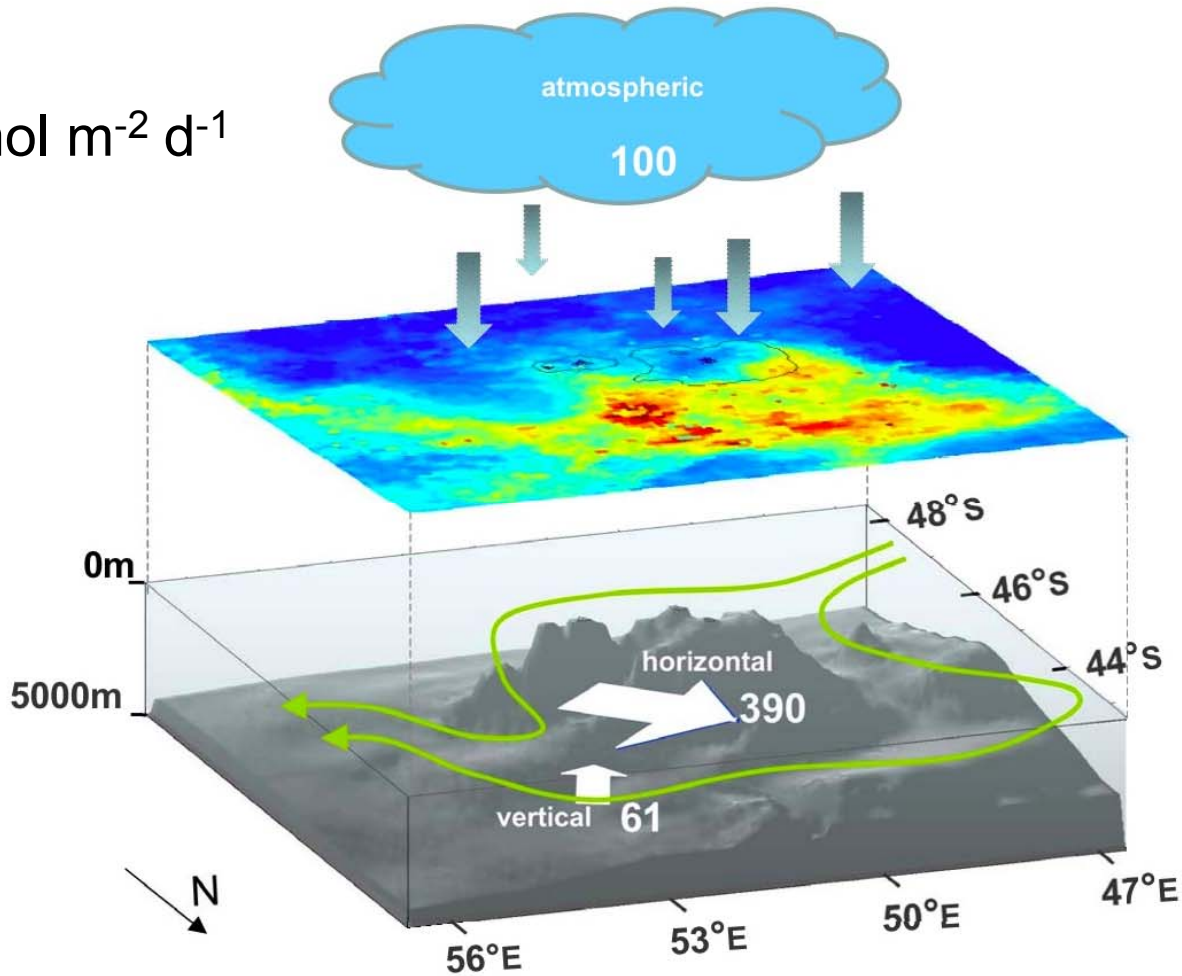
General Hypothesis: HNLC water sweeps north across the Crozet plateau and becomes *naturally* Fe fertilised from shallow (<1000m) sediments

# Control of phytoplankton bloom and POC/N export around the Crozet Islands



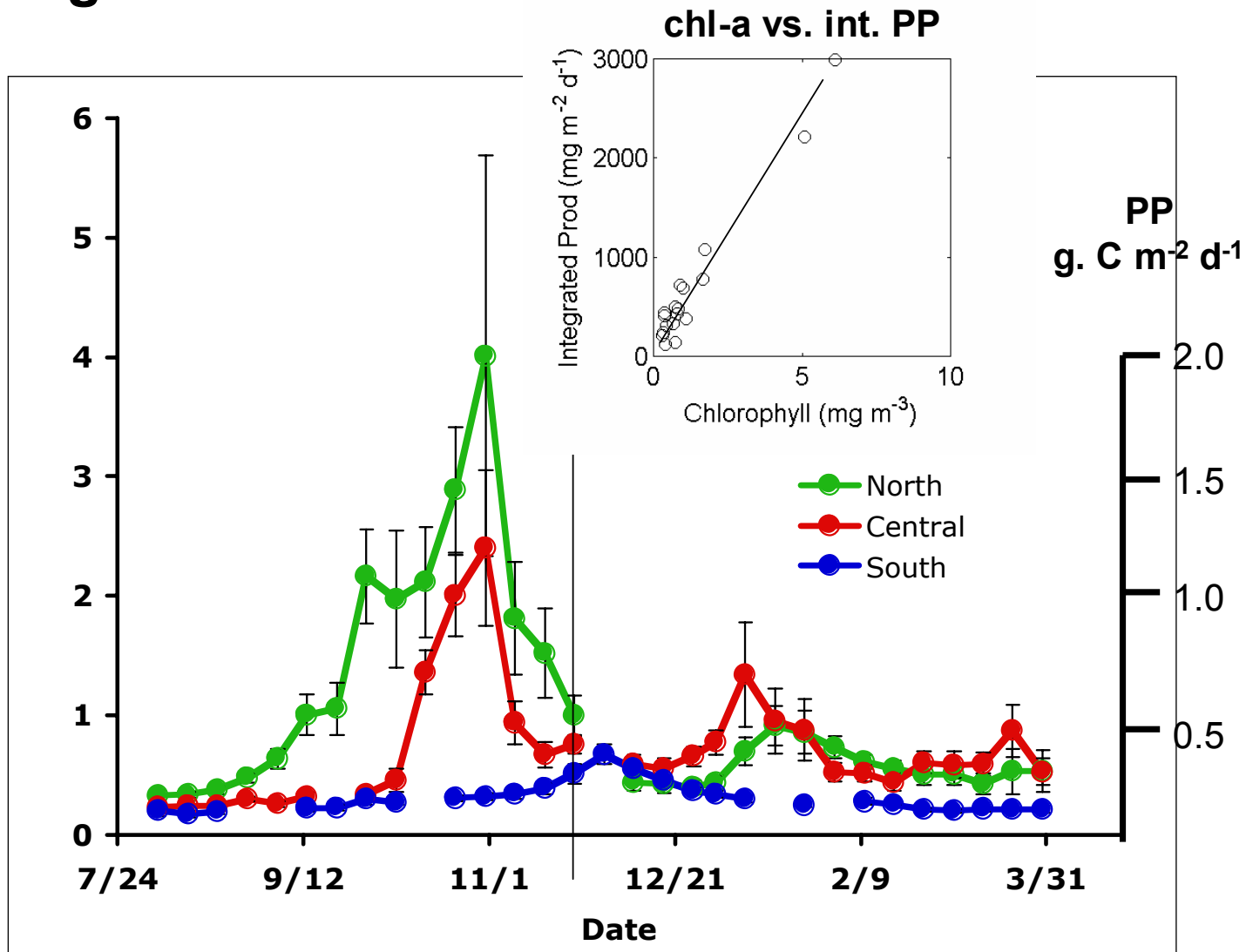
# 1) Iron inputs: DFe fluxes ( $\text{nmol m}^{-2} \text{d}^{-1}$ )

Total:  
 $551 \text{ nmol m}^{-2} \text{ d}^{-1}$

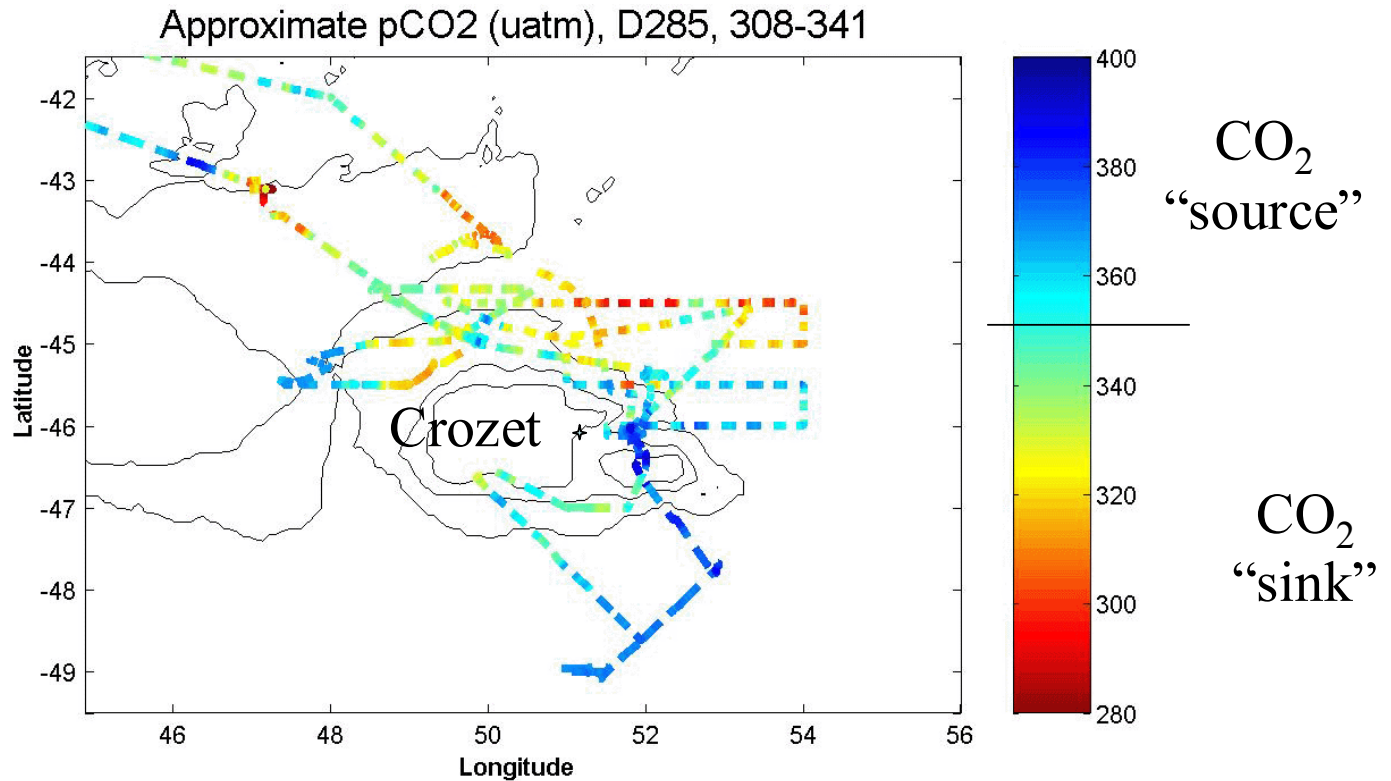


Planquette et al., DSR II, 54: 2007

## 2.) Magnitude of bloom



# Under-saturation of surface pCO<sub>2</sub> in regions of high production



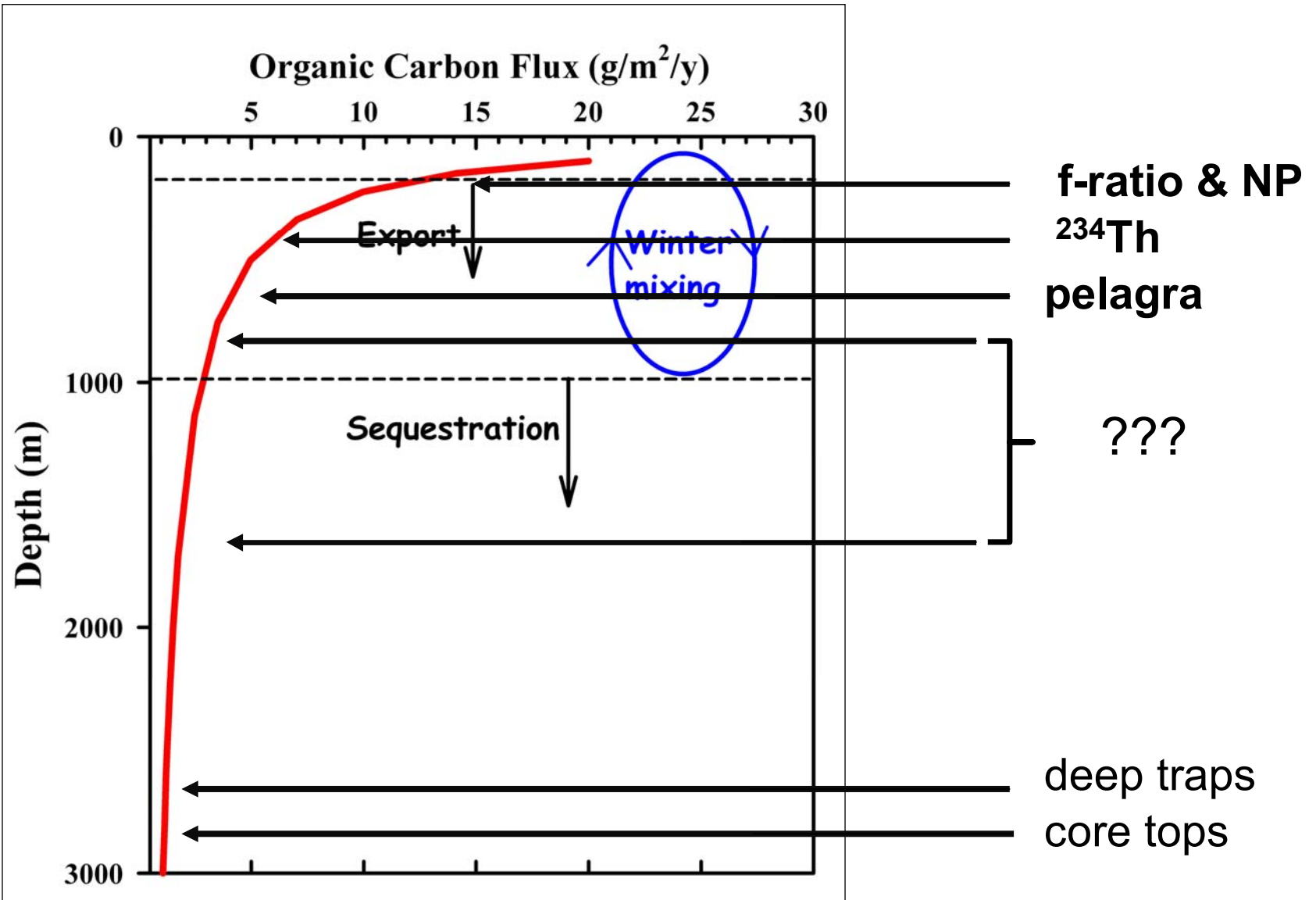
D. Bakker et al, DSR, *in press*

# Carbon Export Measurements

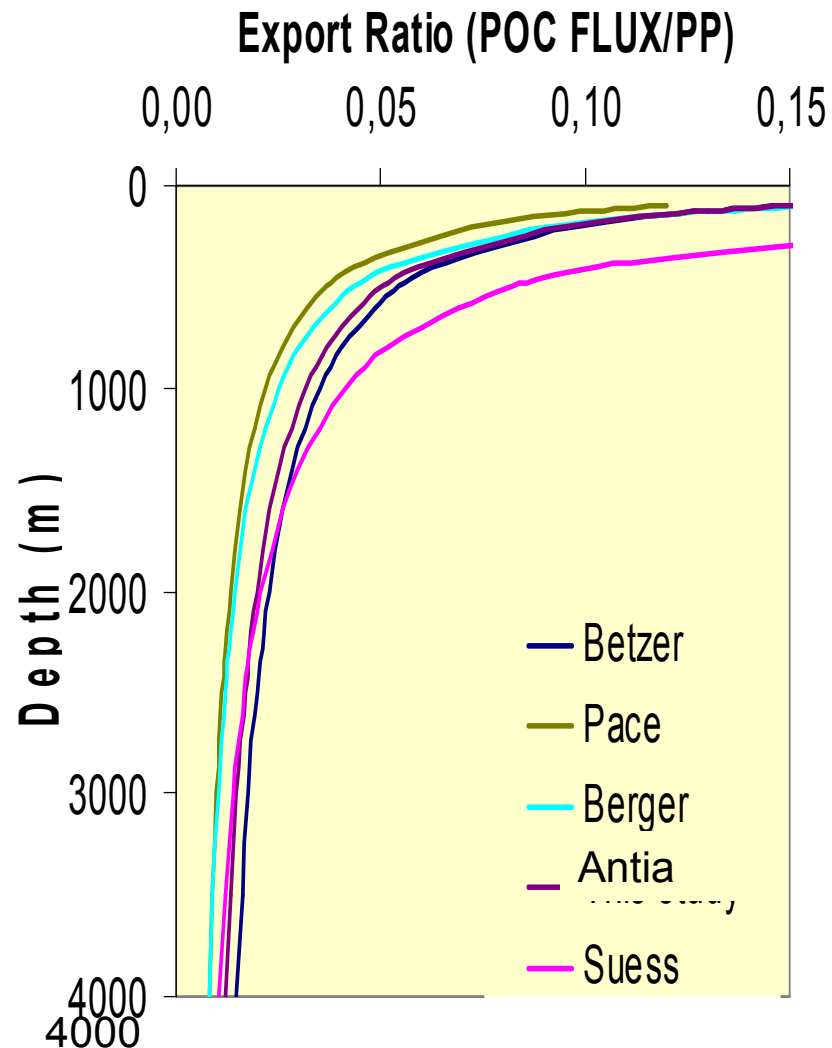


- 1)  $^{15}\text{N}$  stable isotopes & f-ratio. (~55m) (Lucas)
- 2) “Pelagra”; Pelagic lagrangian sediment trap @ ~250m. (Sanders)
- 3)  $^{234}\text{Th}$  /  $^{238}\text{U}$  (~150m) (Morris & Thomalla)
- 4) long-term (~1yr) deep moorings (~3000m)

# Does + Fe result in C sequestration?



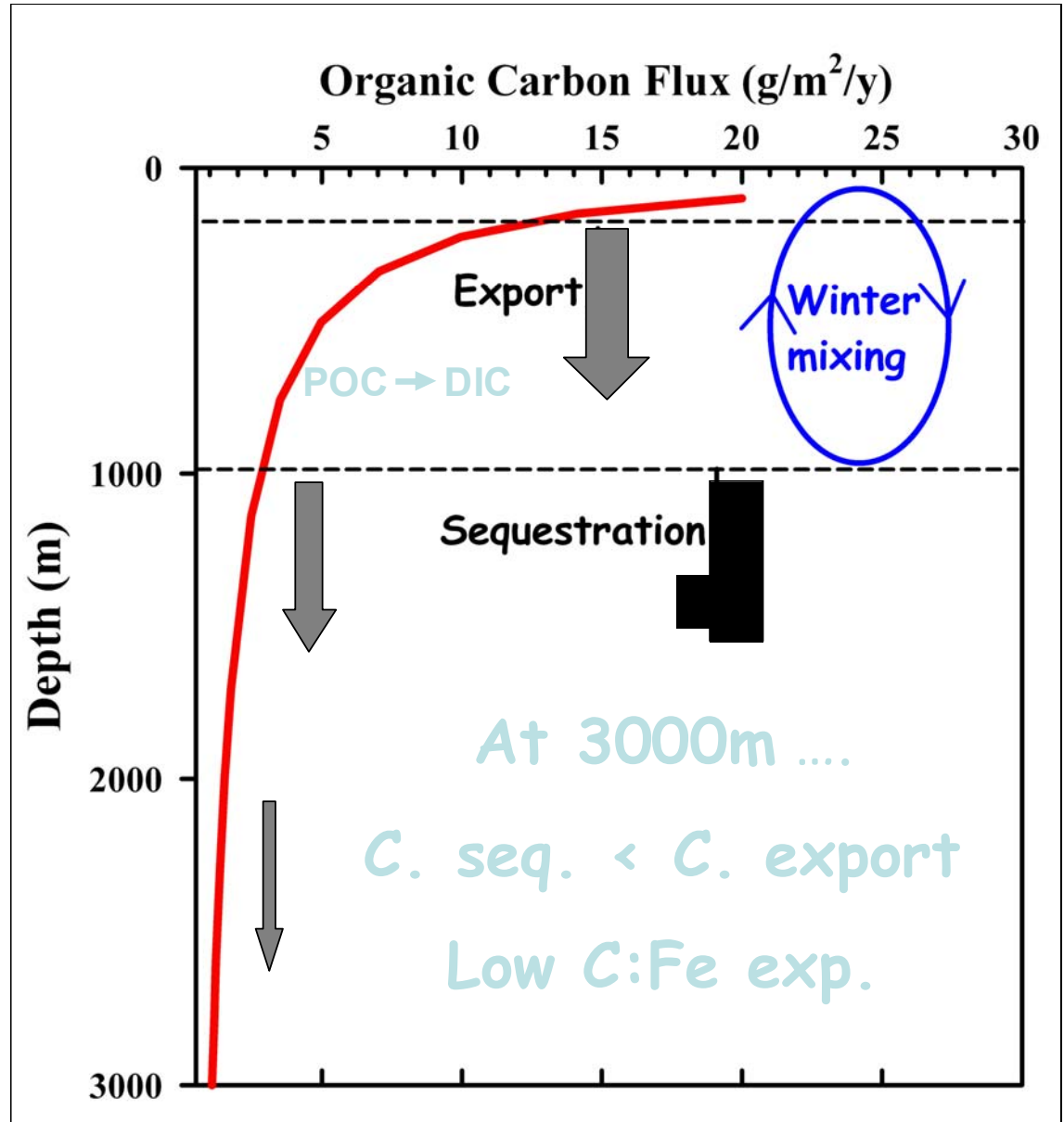
# Carbon export results consistent with Martin curve



# Limited C sequestered Deep Traps



3000 m



# Reality of C. removed by Fe.

Interglacial to Glacial      ~60 billion tonnes

Early 1990's climate models based on lab exps:-

Fe-fertilisation of entire S.O. removes 1-2 billion tonnes C / yr  
(~10-25% of current CO<sub>2</sub> emissions)

Natural Fe fertilisation (Blain et al., '07, Pollard et al., '07:-

1 tonne Fe exports just ~1000 tonnes carbon to 3000m

(Scaled to entire S.O. is << 1% of current CO<sub>2</sub> emissions)

# Benthic Communities?

Comparisons:

M5 - downstream +Fe region

M6 - HNLC -Fe region

**+ Fe**



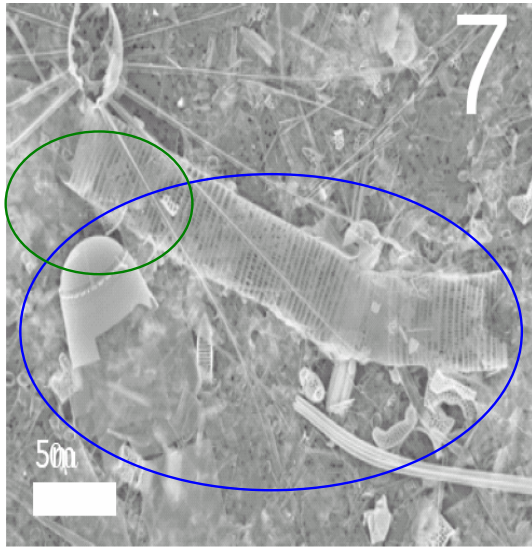
**M5**

**- Fe**

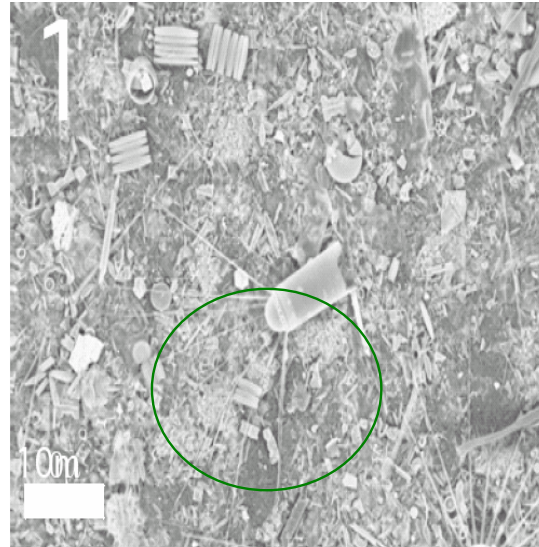


**M6**

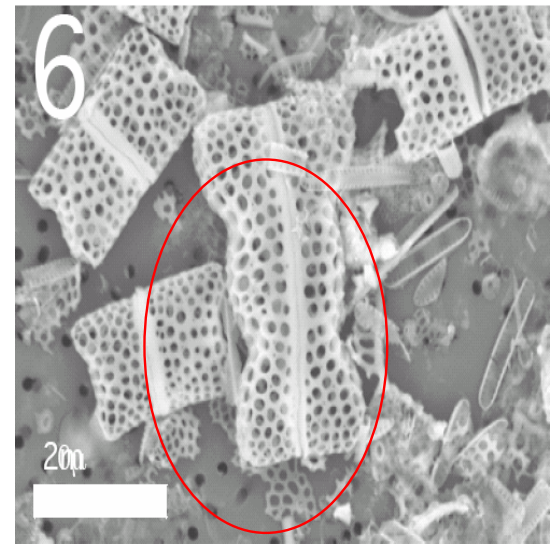
# Particulate material rich in diatom frustules



*D. antarcticus*

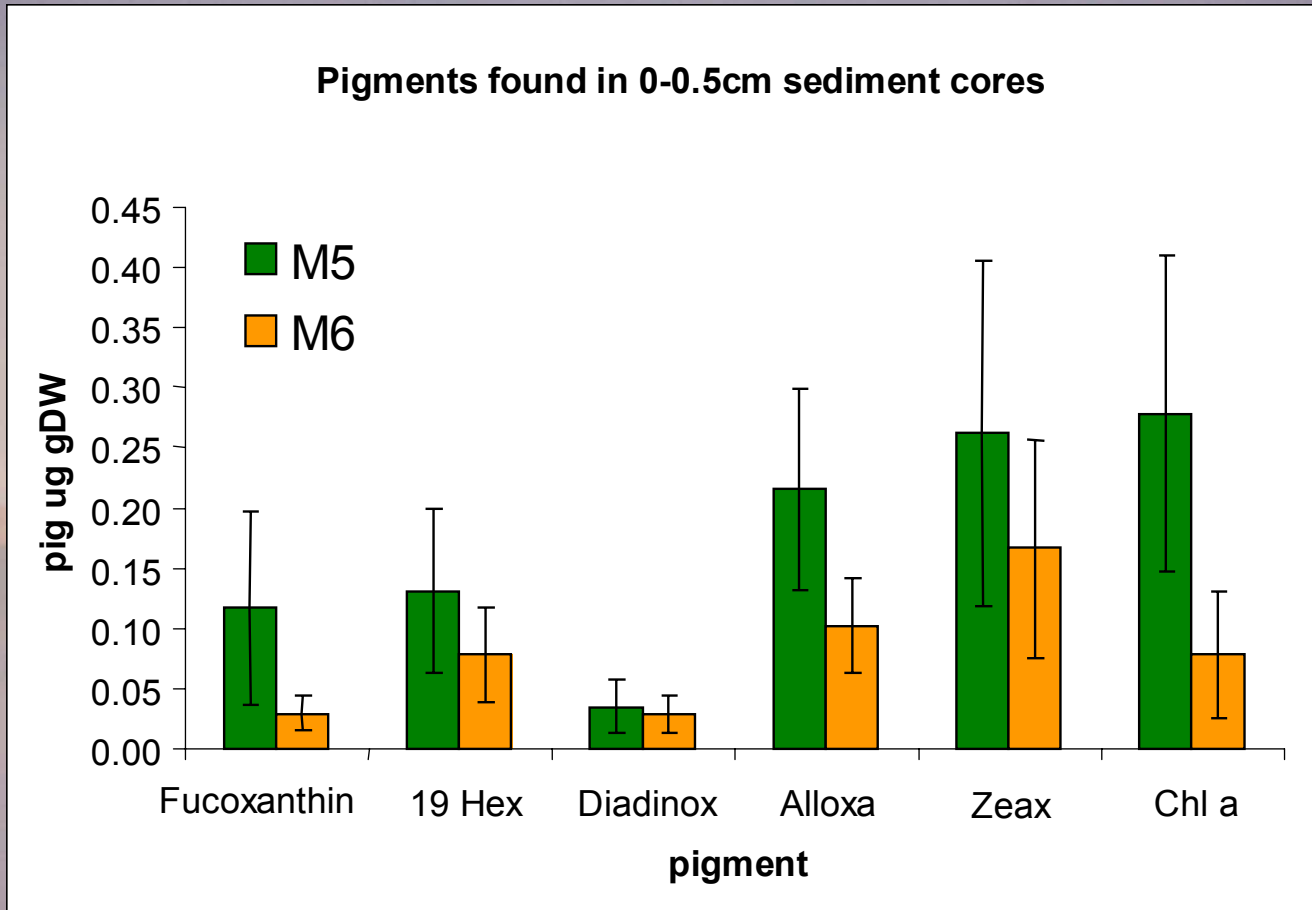


*C. criophilum*



*E. antarctica*

# HPLC Pigment analyses

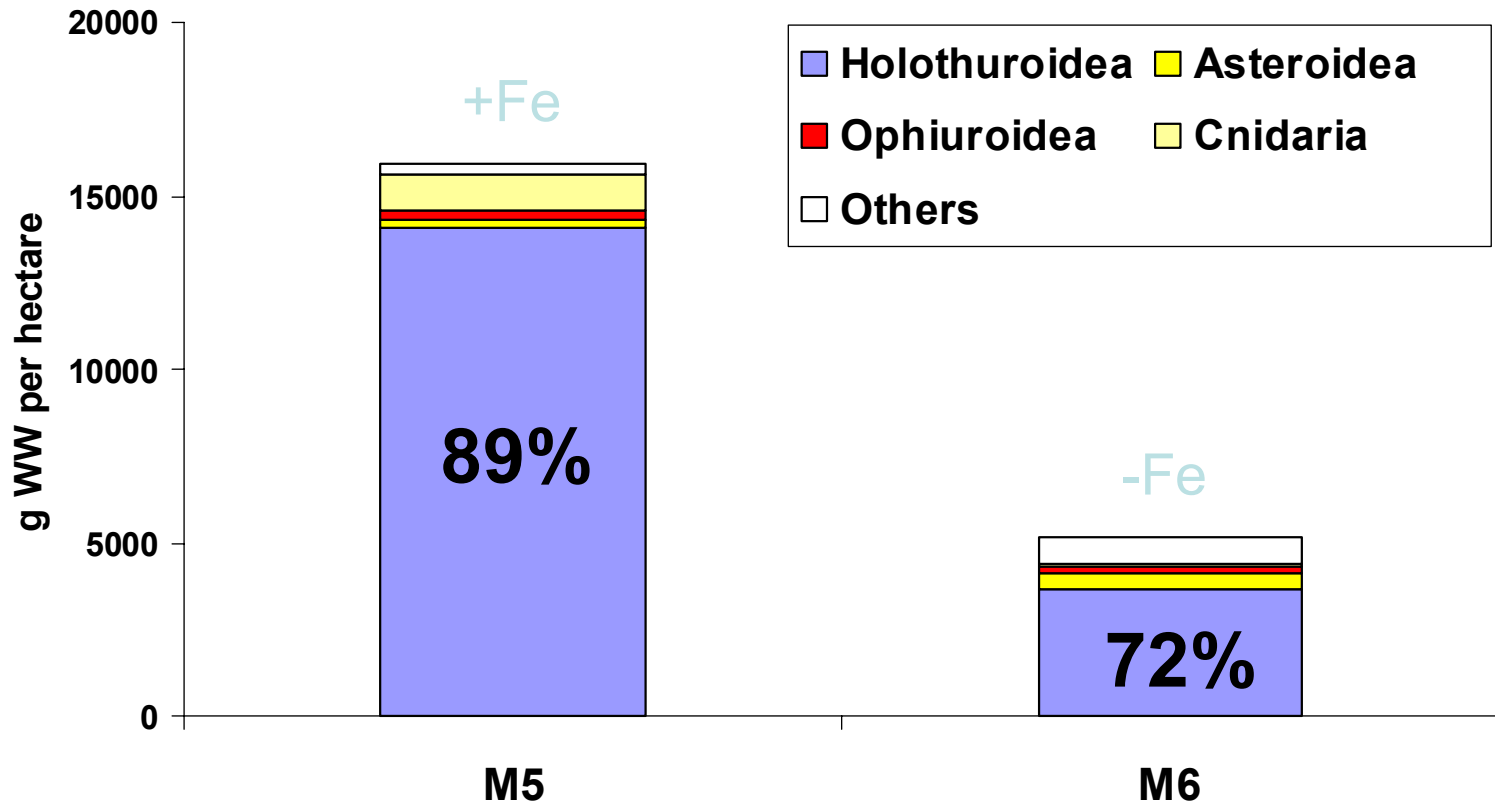


**Stations significantly different ( $P < 0.02$ )**

**Tania Smith**

# Mega-benthos Biomass

## Megafauna Fresh WW Biomass





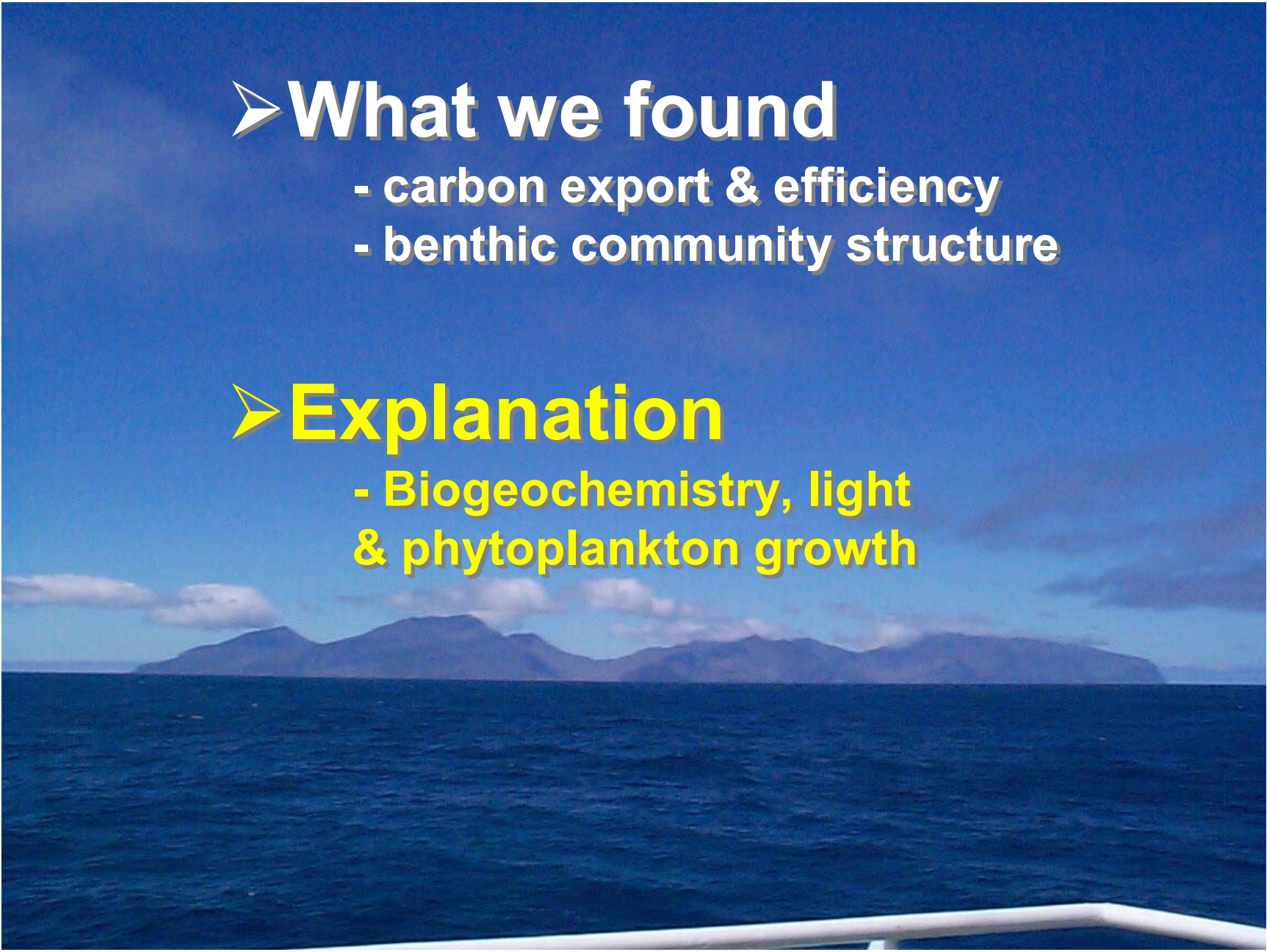


## ➤ **What we found**

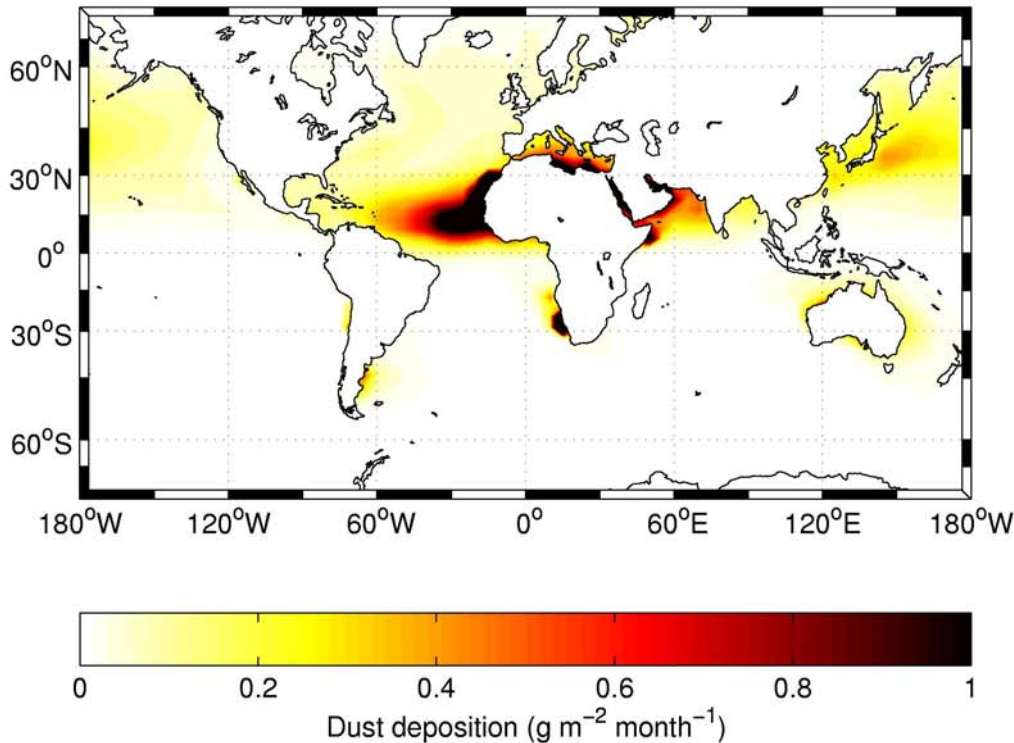
- carbon export & efficiency
- benthic community structure

## ➤ **Explanation**

- Biogeochemistry, light & phytoplankton growth



# Atmospheric Dust Deposition ( $\text{g m}^{-2} \text{ month}^{-1}$ )

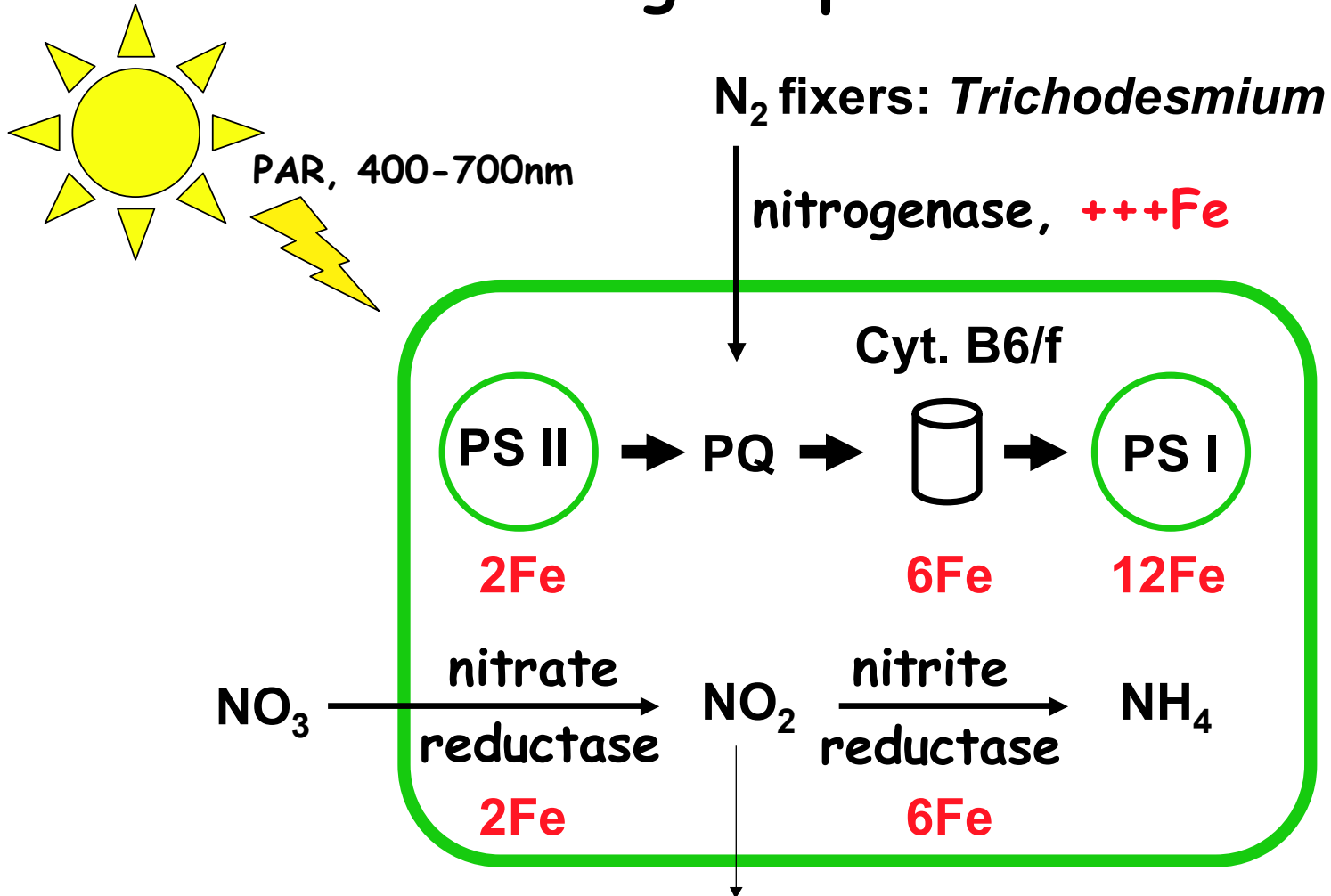


**Little / no Fe  
input into S.O.  
or Equ. Sth.  
Pacific / Atl.**

**Results in low  
biomass & PP  
dominated by non-  
diatoms & small  
cells**

Iron limitation has now been unequivocally established,  
Martin et al. 1994; Coale et al. 1996, 2004;  
Boyd et al. 2000, Tsuda et al. 2003.

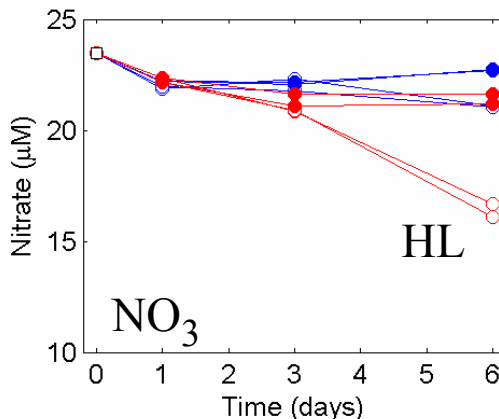
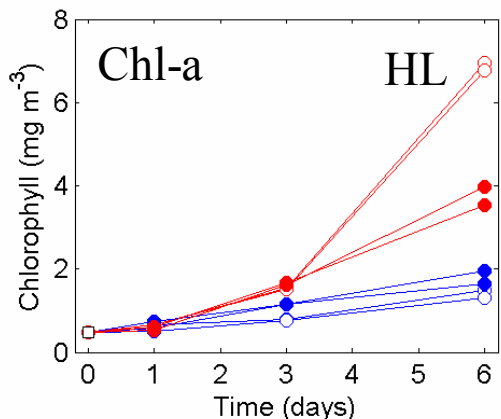
# Light & Fe dependency of Photosynthesis & Nitrogen uptake



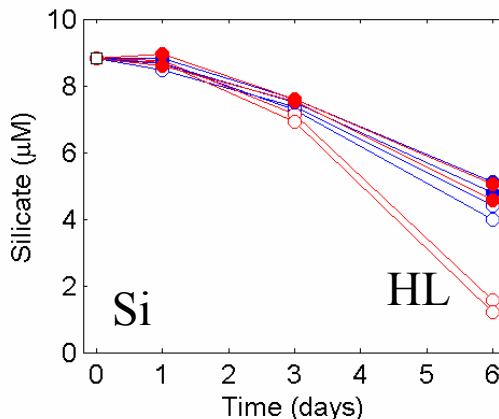
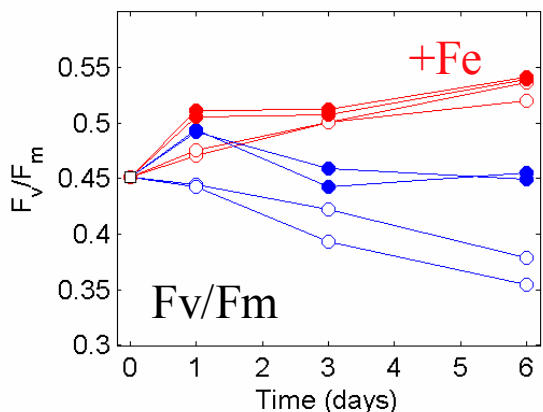
# Bottle incubations: M3 Diatom Dominated

Red: +2nM Fe, Blue: controls

● Low light ○ High light



Light + Fe control  
NO<sub>3</sub> uptake.



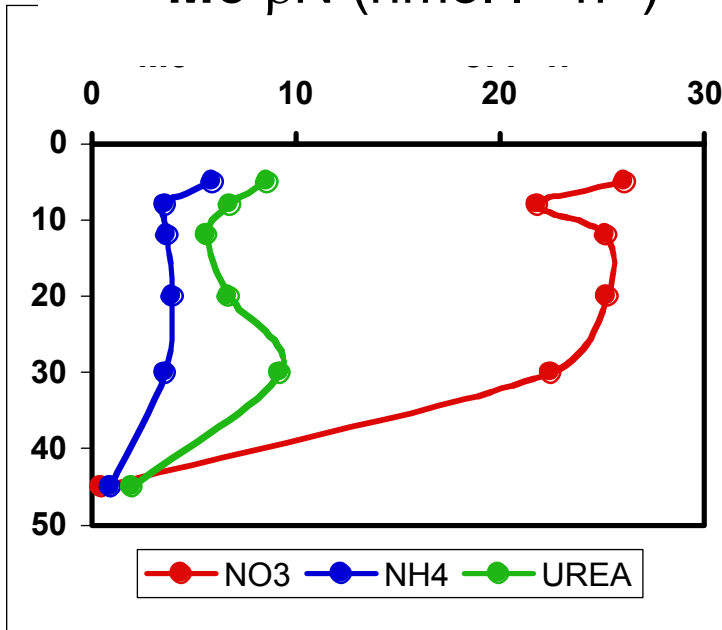
Si uptake is  
relatively  
insensitive to both  
Fe and light.

Physiological “health”  
(Fv/Fm) improves

# Fe impacts Nitrogen uptake rates

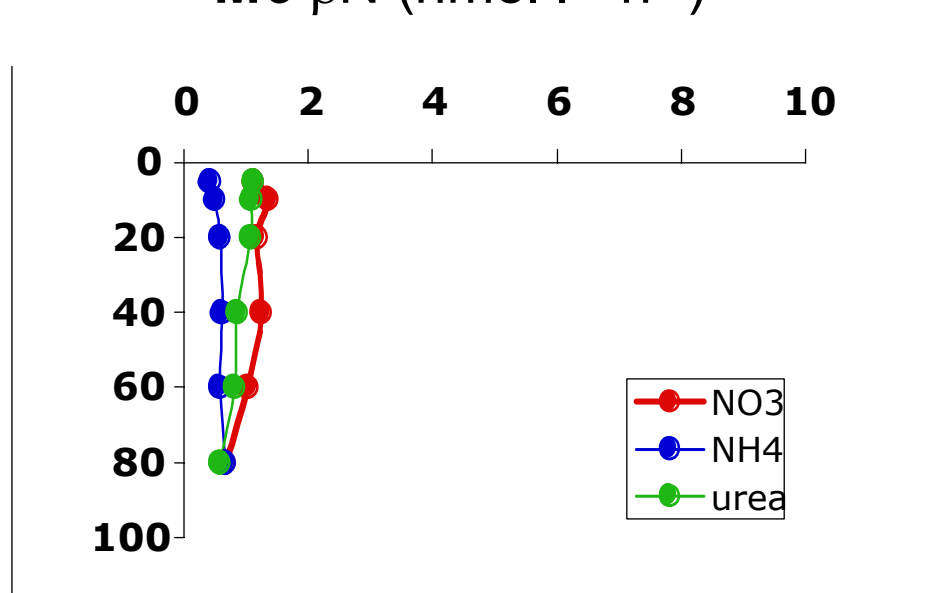
Lucas et al., DSR, 54: 2007

### M3 $\rho N$ ( $\text{nmol l}^{-1} \text{h}^{-1}$ )



North: Bloom  
(+Fe)

### M6 $\rho N$ ( $\text{nmol l}^{-1} \text{h}^{-1}$ )



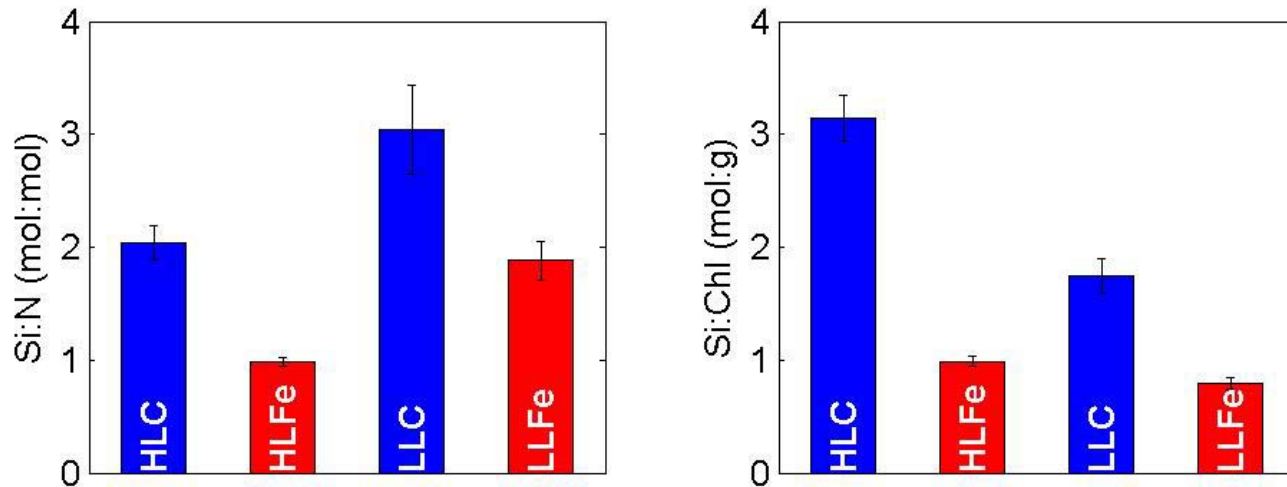
South: HNLC Control  
(-Fe)

# Low Fe impacts on $\rho\text{NO}_3$ and chl-a synthesis

	<b>PON/chl-a</b> ( $\mu\text{mol}:\mu\text{g}$ )	<b><math>\rho\text{NO}_3/\text{PON}</math></b> ( $\text{VNO}_3 \text{ d}^{-1}$ )
<b>North</b> (Post-bloom)	<b><math>2.3 \pm 0.3</math></b>	<b><math>0.05 \pm 0.02</math></b>
<b>M3 Bloom</b> (+Fe)	<b><math>0.98 \pm 0.5</math></b>	<b><math>0.10 \pm 0.03</math></b>
<b>South</b> (HNLC)	<b><math>3.23 \pm 0.8</math></b>	<b><math>0.02 \pm 0.01</math></b>

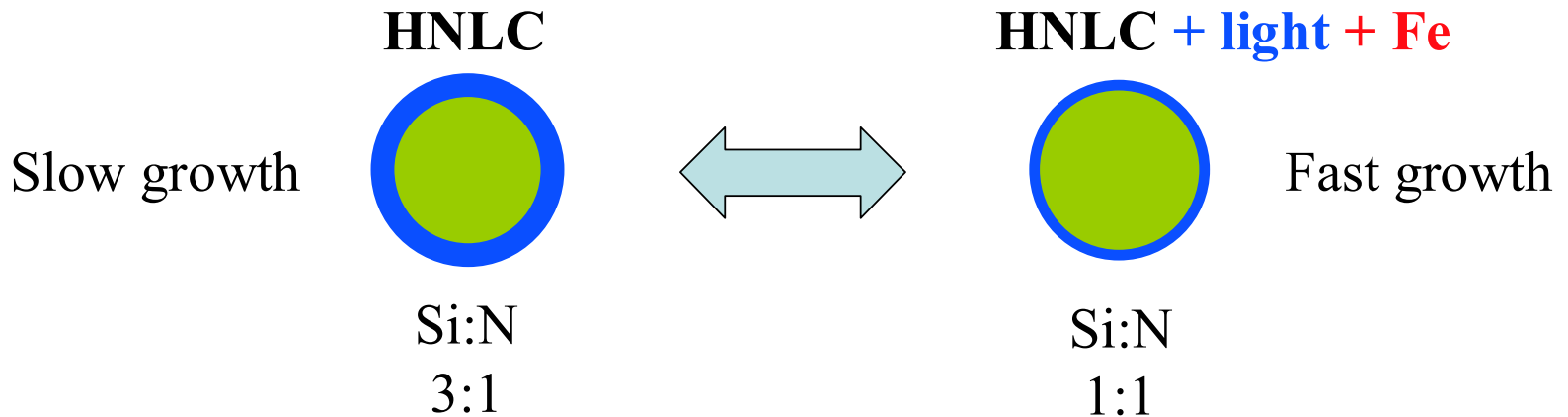
# Fe & the Global Si Cycle

Iron and light drive variability in Si:N draw-down



Moore et al., DSR II, 54: 2007

# Un-couples Si:N Redfield Ratio (~1:0.9) and changes Si:N draw-down ratio



## Implications for-

Glacial “Si leakage” hypothesis

“Rain-rate” of biogenic silica

Ballasting / sinking rate of diatoms

Resistance to herbivory by zooplankton

## **Key Conclusions (1)**

**Annual integrated POC flux is 5 x higher in +Fe region**

**Sequestr. efficiency (POC flux : Fe added) is 700 mol:mol (3000m)**

**Artificial Fe addition exp., efficiency is ~7000 mol:mol (100-200m)**

**POC flux is depth & Fe sensitive**

**Benthic community responds to Fe mediated bloom & C. export**



## Key Conclusions (2)

+ Fe allows nitrate to be used since PSII, PSI & nitrate reductase are Fe dependent

LGM  $p\text{CO}_2$  decrease was ~80-100 ppm

Fe fertilisation to SO contributes just 30-50 ppm

But “Si leakage” to lower latitudes could account for balance.

But - our evidence suggests NO silicate leakage.

Other mechanisms to account for  $p\text{CO}_2$  decrease in LGM?

Thank you!